

Water Resistivity for Western Latvia Ordovician Zone  
Otis P. Armstrong, P.E.

Abstract:

These considerations are given because of uncertainties in Soviet era well logs. Specifically, a lack of water analysis from which Ordovician water resistivity values could be determined. Water Resistivity is an important factor used to calculate porosity and water saturation. Water samples were available for Cambium age, Figure 1, and lower Devonian age, ( 4 to 8 g/l) but missing for Ordovician and Silurian age rocks. A lack of free flowing Ordovician water, was reported to prevented sampling and analysis. Also, the self potential, SP, data were reported to be less than analytical quality. Compounding the problems of SP data were a lack of mud filtrate resistivity values and uncertainty in chart factors.

However, a major shift in SP and deep resistivity,  $R_L$ , appeared across the Ordovician and Cambium boundary, in most well logs; while the Normal resistivity Sonde value,  $R_N$ , remained fairly constant. Also, when making Pickett type plots of Neutron-Gamma Ray deflection vs. log of deep resistivity, there was an improvement in “goodness of fit,  $R^2$ ” for the baseline “ $S_w = 1$ ”, if separate plots were made for Ordovician and Cambium sections. Plus the calculated Pickett values of  $R_w$  for Ordovician averaged higher than  $R_w$  of Cambium sections.

But, a simple analysis of core porosity contradicted the values obtained from the Pickett Plots, SP and sectional resistivity data. This contradiction was resolved by using shale analysis. Details of his novel method are provided in the appendix.

In summary, a water resistivity of 0.13 o-m at 68F, 20C, or 56,000 ppm, NaCl Eqiv., is proposed for Ordovician water in this area. This compares to approximately 120g/l or 120,000 ppm for the mid Cambium section ..

Also, this paper presents a method to ascertain the effect of shale in a formation on resistivity derived water saturation values. This was done by introduction of a shale factor,  $F_s$ , into Archie’s water saturation equation. The Archie equation is revised to read:

$$S_w = \sqrt{(F_s F R_w / R_t)} \quad F_s = \{5C^2\} / [1 + 2C(F R_w / S_w) * (V/R)_s ]$$
$$S_{x0} = \sqrt{(F_{s-x0} F R_{mf} / R_{x0})} \quad F_{s-x0} = \{5C^2\} / [1 + 2C(F R_{mf} / S_{x0}) * (V/R)_s ]$$

With these equations water saturation can be rapidly interpolated either directly or with aid from a Picket plot, as detailed in the appendix. The primary use here is for plotting effect of shale in Pickett chart.

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- 1.3 Resistivity Ratio between Cambrium and Ordovician Sections
- 1.4 Ordovician Water Mineralization by SP ratio
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# Water Resistivity for Western Latvia Ordovician Zone

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## 1.1 Overview

The top of the Western Latvia sequence is typically thought to range from Silurian to Devonian with a cover of 20 to 60 m of glacial quaternary. The top of Western Latvia Ordovician lies at a depth of between -825 m below sea level to -1300 m, with a thickness of approximately 200m. Average surface height ranges from sea level to +80m above mean sea level. The Ordovician out crops on an Estonian island, Saaremaa, about 150 km N-NE of the area of interest. The Ordovician top to the west of Latvia on the Swedish isle of Gotland is approximately ½ the depth that found in western Latvia.

Typically, a few meters of blue clay overlies the Cambrium top in this area of Latvia. Below the bottom Ordovician clay boundary is often found a highly cemented glauconitic sandstone, followed by the rest of the Cambrium sand and clay sequence of 100 to 200m thickness. At least one paleontology writer theorized an ancient continent of Baltica in a more temperate zone than other basins during development of Cambrium and Ordovician. This less tropical zone is given as the cause for modest thickness in Ordovician zone of the area.

The Latvia substructure was drilled in Soviet era by over 100 wells, with extensive coring and testing. In all Soviet work, there are no reports on water mineralization for the Ordovician. Perhaps, this was due to a lack of water flow in the Ordovician. It is hoped, the following considerations resolve the uncertainty of water mineralization values for Ordovician...

## 1.2 Cambrium Water Mineralization

The basis for most calculation is water mineralization analysis for the Cambrium section. The effect of divalent ions in solution did not alter water resistivity by more than 3% from values obtained by NaCl tables. For this work, all resistivity values

Location, Upper. Cambrium	.mg/l	T: C, F	R, o-m
Edole, Kuldiga, Vergales	105	22, 72.0	0.073 o-m
Durbes and Astere:	112.5	24, 74.3	0.068 o-m
Bernati	120	40, 104	0.047 o-m

are from NaCl tables. Figure 1 shows some Cambrium water resistivity results for 9 wells in the area. The resistivity ranges between 0.063 and 0.095 o-m at 75F. The accepted values for upper Cambrium are:

## 1.3 Resistivity Ratio between Cambrium and Ordovician Sections

Table 1 gives a summary of normal and long resistivity values for 14 wells in Western Latvia. Looking at the average results, the ratio of deep resistivity,  $R_L$ , values for the two sections is 5.6 and the shallow resistivity,  $R_N$ , ratio is 1.98. Typically  $R_N$  is related to  $F \cdot R_{mf}$ , Formation Factor,  $F$  and mud filtrate resistivity. Since mud filtrate mineralization is considered constant along the well bore, the ratio between  $R_N$  in Cambrium and Ordovician sections represents variations between  $F$  and temperature gradients (0.9C/100m) in the two sections. Division of the  $R_L$  ratio by the  $R_N$  ratio takes out this effect, giving a ratio of 5.6/1.98 or 2.8. By this method Ordovician water resistivity is  $2.8 \cdot 0.065$  or 0.18 o-m.

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The same result would be obtained if  $R_L$  ratio were compared between sections for which  $R_N$  in Cambrium and Ordovician horizons were equal. In which case the ratio between  $R_L$  for these two sections would be the ratio between water mineralization in Cambrium and Ordovician sections of equal porosity.

Table 1 Average Section Resistivity

	Ord	Ord	Ord	Camb	Camb	Camb
well	RL	RN	Rn/RL	RL	RN	Rn/RL
V50	5.0	20.0	4.0	1.0	6.0	6.0
KAz6d	20.0	15.0	0.8	0.7	5.0	7.1
K5	4.0	3.5	0.9	0.9	7.0	7.8
K4	4.0	4.0	1.0	1.5	5.0	3.3
K3	4.0	4.0	1.0	1.0	5.0	5.0
K13	13.0	70.0	5.4	7.0	20.0	2.9
Ed60	10.0	10.0	1.0	0.9	5.0	5.6
Ed17	14.0	12.0	0.9	1.0	7.0	7.0
D35	8.0	7.0	0.9	3.0	5.0	1.7
D15	16.0	11.0	0.7	1.2	6.0	5.0
D14	16.0	14.0	0.9	1.5	7.0	4.7
D13	5.0	3.0	0.6	0.6	4.0	6.7
B6	15.0	4.0	0.3	2.0	4.0	2.0
B20	8.0	8.0	1.0	3.0	8.0	2.7
avg	10.1	13.3	1.4	1.8	6.7	4.8

The ratio of  $R_N/R_L$ , for a given section, typically eliminates formation factors, and the result should be  $R_{mf}/R_w$ . Since  $R_{mf}$  should be the same along the borehole, taking the ratio of  $R_{mf}/R_w$  for Ordovician and Cambrium can give a pseudo water resistivity ratio, which for the below table comes to 2.7, indicating  $R_w$  Ordovician of 0.180-m at an average Cambrium water R of 0.065.

When shale and temperature factors are applied, Eqn.A-25, Ordovician water calculates to be 0.14 o-m, see appendix for details.

$$R_w/R_{mf} = (R_t/R_{x0}) * [1 + 2C(FR_w)*(V/R)_s] / [1 + 2C(FR_{mf})*(V/R)_s] \quad \text{A25}$$

1.4 Ordovician Water Mineralization by SP ratio

The Cambrium in the area of interest, is mostly medium porosity sandstone with intermixed shale, while Ordovician is mostly low porosity carbonates, shale or mixtures. In the case of resistivity, this shift could be influenced by porosity, lithology or changes in both. However, SP data in a thick, HC free bed, is responsive only to changes in shale volume content. So the difference in maximum SP deflections between two clean, thick beds can be related to Ordovician  $R_w$  as follows:

$$(-SP/K) = (\log R_{mf}/R_w) = (\log R_{mf} - \log R_w) \quad \text{Eq1}$$

Since the thermal gradient is low and less than 200m between sections, an average K could be assumed effective and  $R_{mf}$  should also be about the same in the two sections. (if SP deflection is not fully developed, flat lined, then bed thickness correction should be applied;  $SP = SP_{def} * CF_z$ , where  $CF_z$  is always greater than or equal to 1.00)

$$CF_z = \exp[1.245(\ln(R_i/R_m))/z' - 0.591/z' - 0.133], \quad z = \text{bed thickness, ft} \quad \text{Eq.1a}$$

Comparison of maximum SP's between Ordovician and Cambrium gives the following simplistic equation for  $R_w$  in Ordovician:

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$$(R_w)_{ordv.} = a \log(\log(R_w)_{cmb} + [(SP)_{ord} - (SP)_{cmb}] / K + C) \quad \text{Eq. 2}$$

The chart factor, C, is the offset factor which makes  $(-SP/K)$  equal to zero when  $R_w$  equals  $R_{mf}$ .

Results of this analysis for Cambrium  $R_w = 0.065$  is given in Figure 2, below:

	F=	1/φ <sup>2</sup> carbs	0.81/φ <sup>2</sup> consol sands	A35
The statistical average for this method is 0.21, the	Log(φ, v/v)	=	U - V*ND	A38
result for well K13 was discarded, being 3 times greater than the average.	At F =10	$R_w = R_t/10$	$R_{mf} = R_{xo}/10$	A40a, b

1.5 Ordovician Water Mineralization by Pickett Plot

When the Pickett method was applied to Ordovician section of 7 well logs, 0.13 was the average  $R_w$ , with a range of 0.09 to 0.18 o-m. This method made an educated assumption about the porosity range. The Cambrium section average for 4 wells was 0.049 o-m.

The method details are given by the Appendix. Briefly it is as follows :use Eq.A35a or b to calculate porosity for F=10, then apply A38 to get ND, and lastly, apply 40a,b to arrive at  $R_w$  or  $R_{mf}$ .

1.6 Ordovician Water by Shale Parameters

The Appendix proposes to calculate  $R_w$  using shale parameters, those results arrive at an Ordovician water salinity of 0.121 o-m at 68F, 0.091 o-m at 93F. This method does not rely on Cambrium water analysis. In this instance, core porosity was used to determine F and plots the product of F and shale volume,  $V_s$ , against the ratio F over log resistivity, Eq.A16.

The slope and intercept of the regression line provide values of Shale resistivity and aqueous phase resistivity.	$F/R_t = FV_s / (5CR_s) + 1/(R_w 5C^2)$	Eq.A16a
	$F/R_n = FV_s / (5CR_s) + 1/(\{R_{mf}\} 5C^2)$	Eq.A16b

Figures 6 and 7 plus the below tables of Normal and Long survey tools were taken for well Bernati 2, Ordovician zone using core open porosity. The result for RL is given in Table 2 and Figure 6.

Table 2: B2 Plot Results, RL,

The average Ordovician water resistivity calculated by shale equations (column B) is about 0.09, at 40C. This compares to 0.05 (E) when determined by Archie Equation with open porosity, which ignored shale effects.

	Rw	Rs	r <sup>2</sup>	Rw=RL/Fa	Rw=RL/Fe
RL	0.099	1.61	0.38	Shale eqn	Regr plot
RL avg	0.091	1.75	na	0.052	0.061
A	B	C	D	E	F

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When porosity is increased by  $(1+Vs)$ ,  $R_w$  calculates higher, 0.06, (F) when ignoring shale effect,  $F_s=1$ . Shale resistivity, is determined from regression equation intercept, Eqn. 14 a & b. However, since shale conductivity theory implies  $R_s$  is independent of saturating water solution, then  $R_s$  is averaged between the values obtained in Figures 6 & 7 for  $R_n$  and  $R_L$ , results in bottom row of Tables 2&3. Using the average of shale resistivity for Normal and Long Sonde gives an average  $R_w$  and  $R_{mf}$  (mud filtrate), bottom row, Tables 2 & 3.

Shale resistivity value,  $R_s$ , for Ordovician of well B2, when corrected for temperature to 68F is 2.3 o-m. This value is very close to the statistical value of 2.8 for the wells listed in Table 4.

The water resistivity value  $R_w$  for Ordovician of well B2, when corrected for temperature to 68F is 0.12 o-m. This value is close to the statistical average value of 0.17 for the wells listed in Table 4, and much higher than Cambrium water resistivity.

Table 3: B2 Plot Results,  $R_n$

	$R_{mf}$	$R_s$	$r^2$	$R_{mf}=R_n/F_o$	$R_{mf}=R_n/F_e$
$R_n$	0.17	1.87	0.55		
$R_n$ avg	0.203	1.75	na	0.071	0.085

Table 3 provides the results for the same calculations for the normal survey sonde. The normal sonde has a shorter spacing and investigation depth. The log data did not provide details about mud filtrate resistivity. Mud gravity's were reported to range between 1.1 and 1.28.

Table 4 gives the values of  $R_s$  and  $R_w$  determined from 11 well logs of western Latvia wells. These value were taken by regression using plot of  $F/R$  vs  $F/V_s$  per Eqn A16a/b. Porosity was from NGR by maximum and minimum porosity, Eqn A38. This method, by default, determines an average shale resistivity.

Table 4 Cambium. & Ordovician  $R_w$  &  $R_s$  results from Shale Plots

Well	Cmb $R_w$	Cmb $R_s$	rr	Ord $R_w$	Ord $R_s$	rr
Ed17op	0.03	0.5	.1	0.21	3.1	.5
B2op	0.06	7.6	.5	0.166	2.46	.77
K13op	0.033	0.76	.54	0.28	1.4	.27
K15op				0.14	3.2	.1
K5op	0.063	0.56	0.03	0.195	4.1	0.82
KD6op	0.024	1.65	0.002	0.13	5.3	0.99
Ed60op	0.03	2.14	0.99	0.159	3.1	0.97
B20op	0.057	7.8	0.999	0.135	2.3	0.91
B24op	0.03	3.2	0.99	0.49	1.7	0.94
AP1op	Na	Na	Na	0.12	2.1	0.83
AP2op	Na	Na	Na	0.18	2.1	0.91
St. Avg	<b>0.041</b>	<b>3.0</b>	<b>0.52</b>	<b>0.17</b>	<b>2.81</b>	<b>0.73</b>

### 1.7. Summary

By most methods, (resistivity ratio, with (2) and without shale factors (2b); SP ratio method (3); Pickett Plot, (PP, 4) and  $R_w$  by Shale plots (5&5c), the Ordovician water resistivity was

calculated to be considerably more than 0.065o-m for Cambrium age.

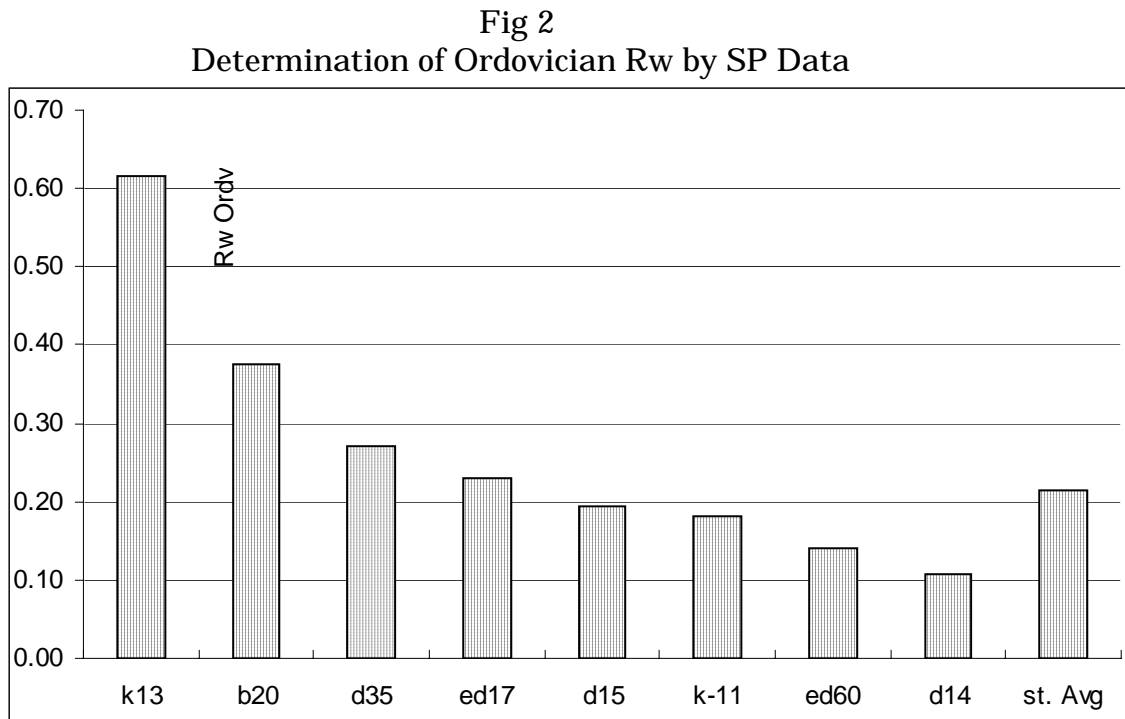
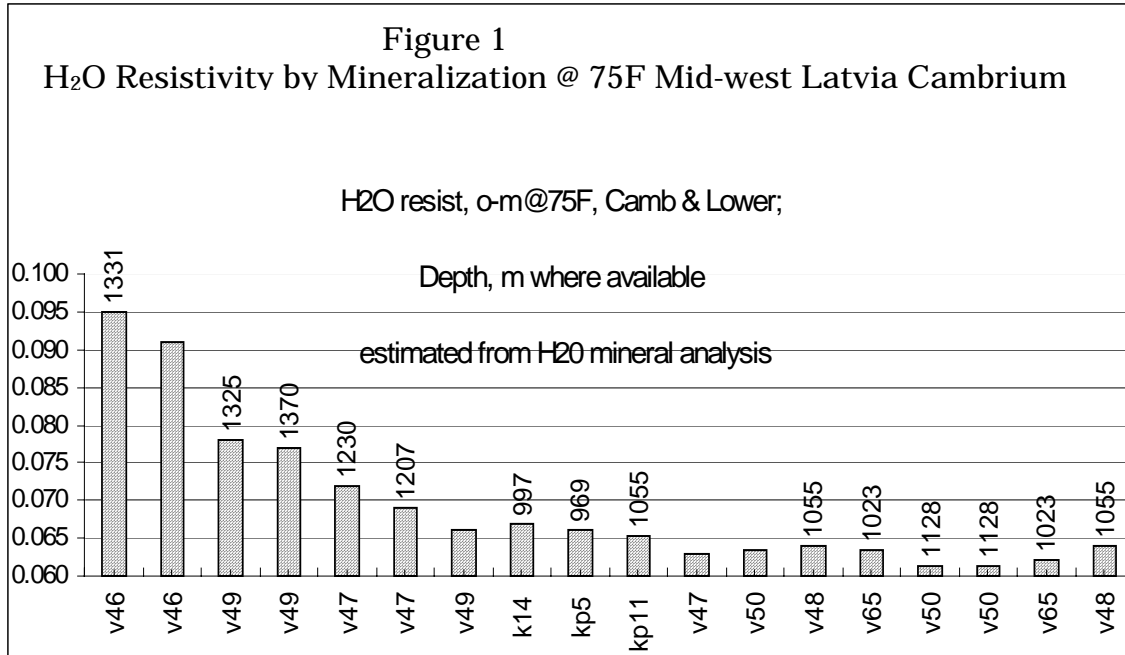
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R-ratio	R-r w/shale	SP	PP	SF core B2	Core B2w/o SF	SF avg	Avg. All
2	2b	3	4	5	5b	5c	#6
Rw	Rw	Rw	Rw-o	Rw-Ov= 20C	Rw-o @40C	Rw-	Rw-o
0.18	0.14	0.21	0.13	0.12	0.065	0.17	0.146

The general average of water resistivity comes to 0.146 for all methods, column 6. The method involving the least amount of assumptions is shale factor plot (5) of porosity data. The shale factor, without core method, (5c) assumes a porosity range to determine F. The SP method (3) required an assumption of mud filtrate resistivity to arrive at the offset factor, C. This assumption of mud filtrate resistivity is marginal, given mud density, log resistivity in large caliber zones, and drilling practice, but nevertheless an assumption. The resistivity ratio method (2) neglects shale factor corrections. The appendix shows that shale corrections are more significant at low porosity or high resistivity. When shale factors are considered, Rw calculates lower (2b). The PP, method (4) assumes a porosity range. This “porosity range” is based on core analysis porosity ranges and the general rule that consolidated rock porosity seldom exceeds 28% and the physical principle that residual porosity is found in all but pure minerals. Shale factors and core porosity values indicate only a remote possibility of pure minerals, thus the porosity range of 6% to 28% for PP analysis. The recommended value is an average of methods, 2b, 4, and 5; which equals 0.130-m at 20C.



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Figure 3  
 Effect of  $R_w$  on need for Shale Analysis  
 Comparison of Wylie Data to Simandoux Equation

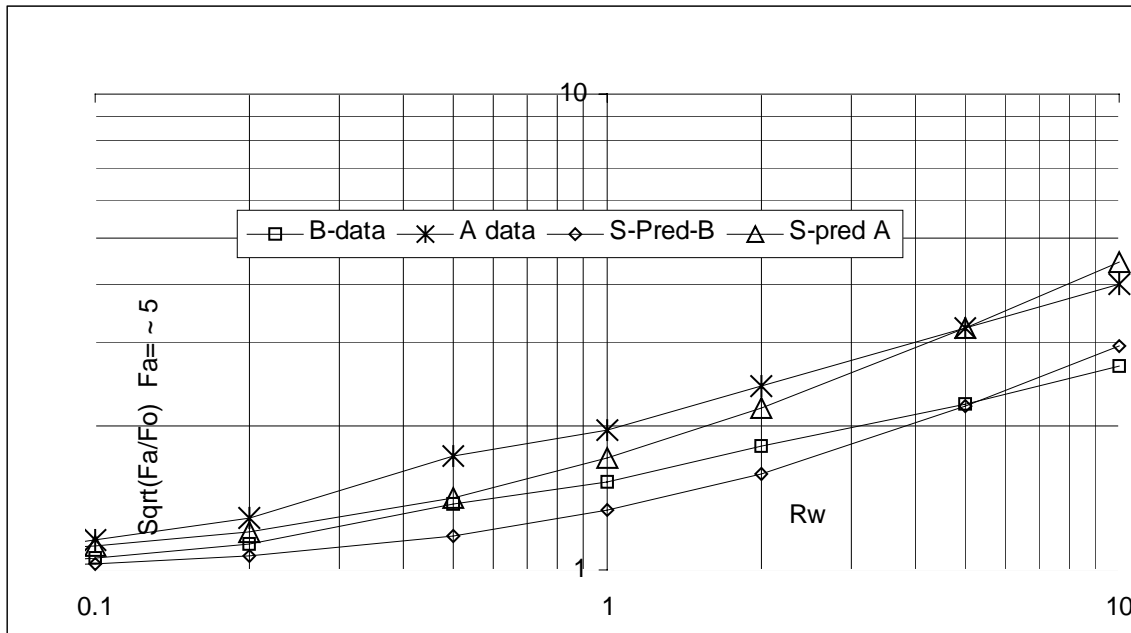
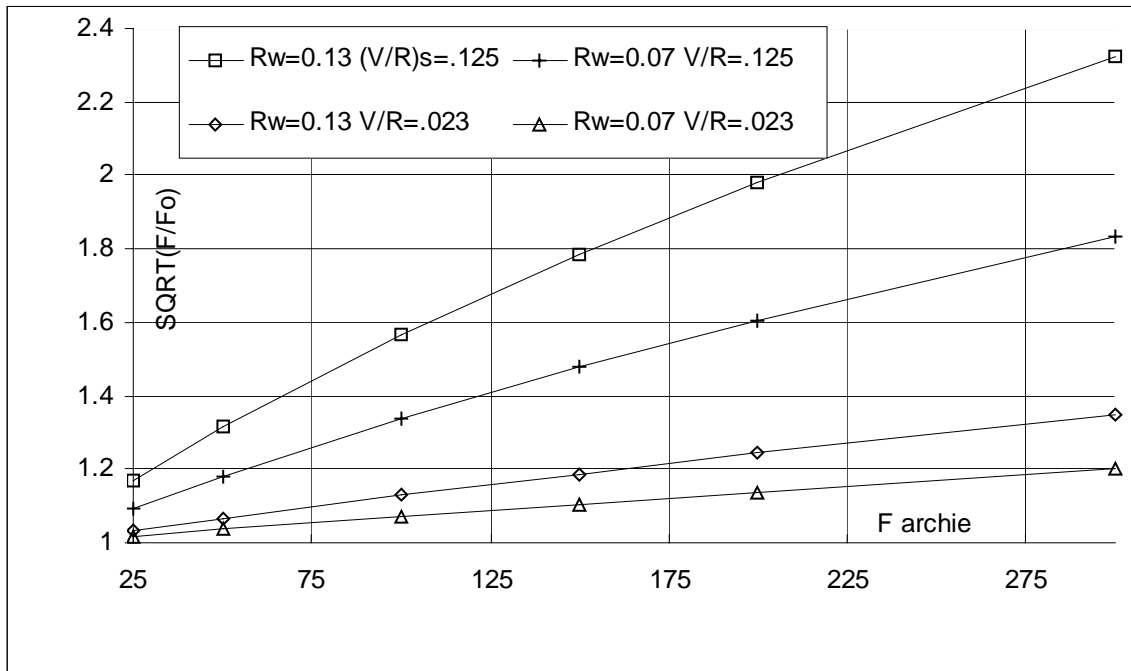


Figure 4  
 Effect of  $R_w$  on need for Shale Analysis at Large F (low porosity)  
 Comparison by Simandoux Equation



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Figure 5  
 Effect of Sw on  $I = R'' @ Sw / R'' @ Sw = 1$  Comparison of Wylie Data at  
 F = 5 and by Simandoux Equation for F=5 and for F = 125, low porosity

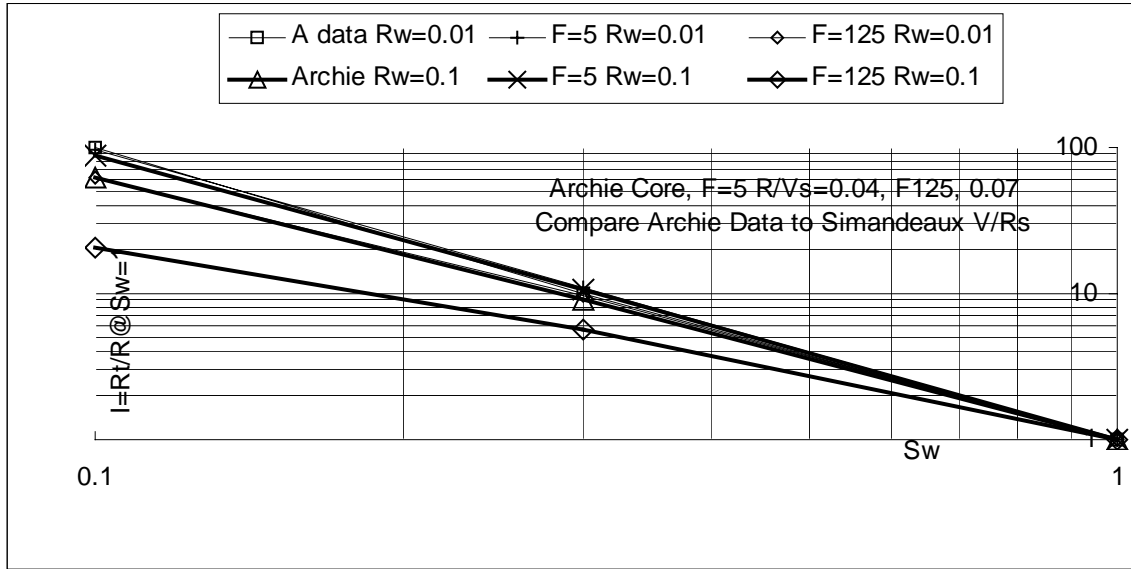
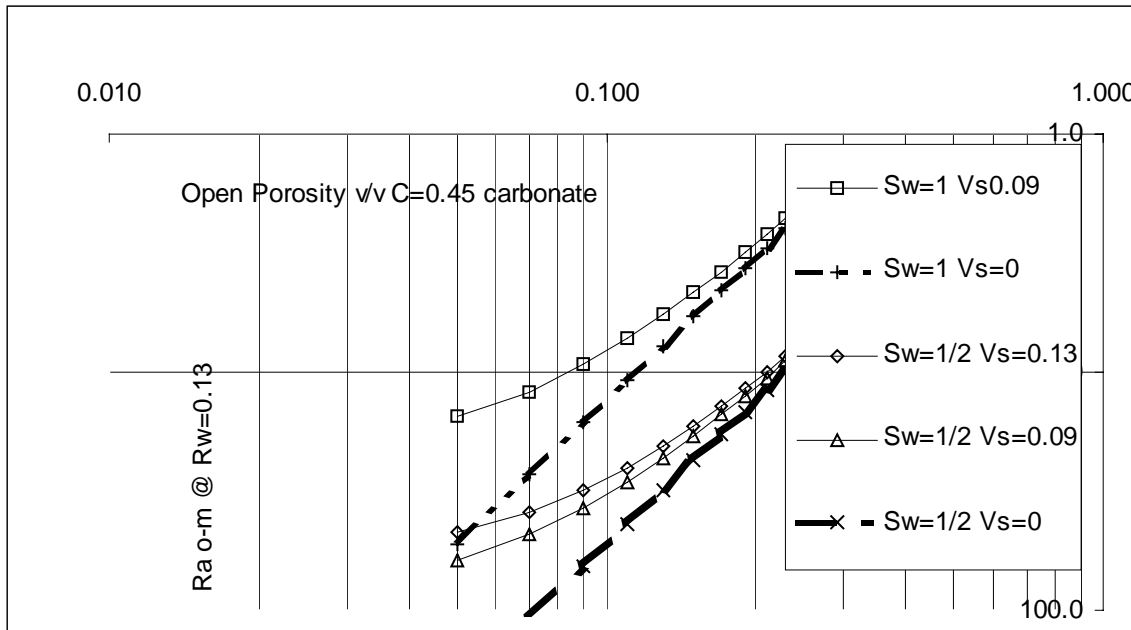
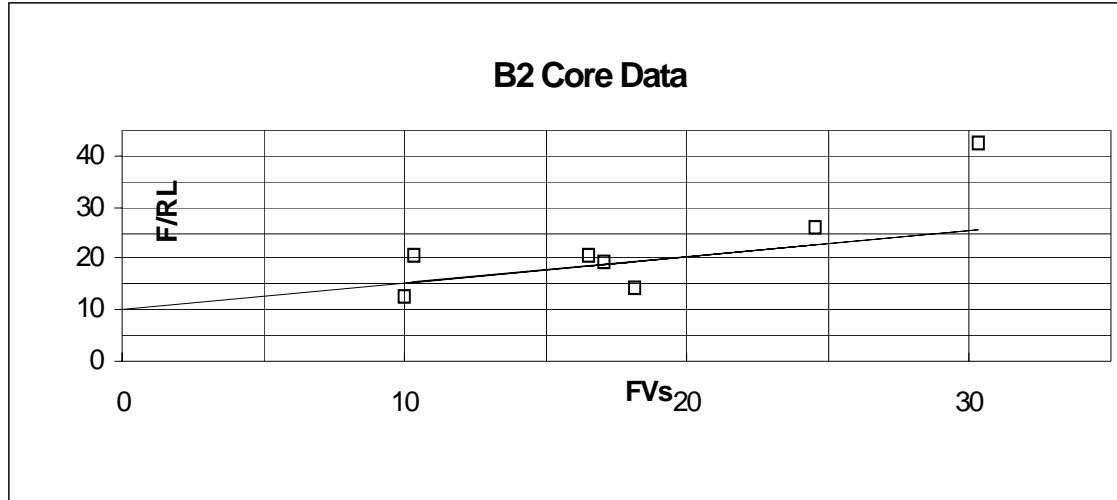


Figure 6  
 Low Open Porosity Effect on Rt for Sw = 50% and 100% at Moderate Shale%

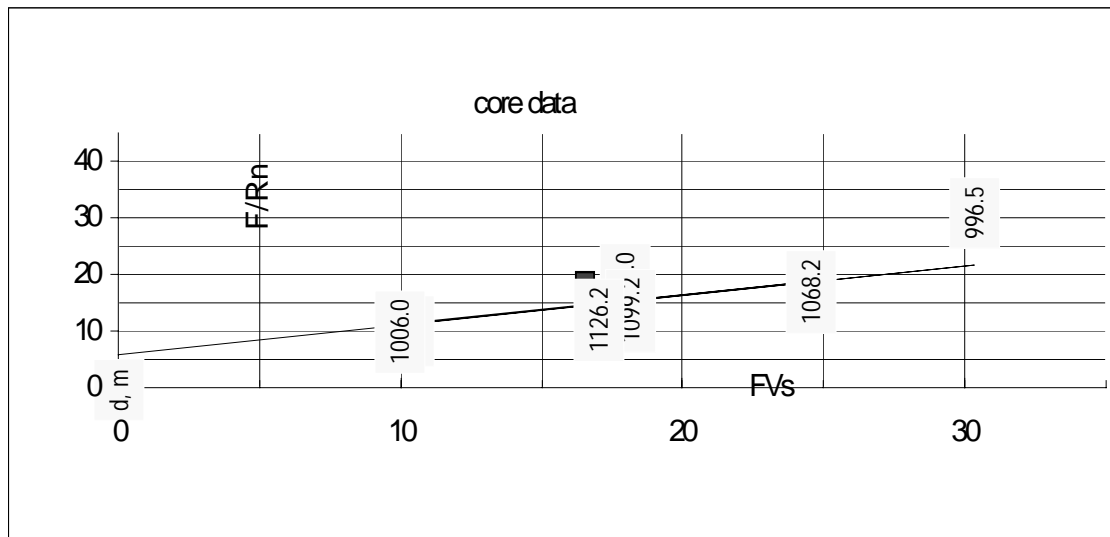


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Figure 7  
 Plot of Long Electrical Survey Tool, RL to Determine Rs and Rw

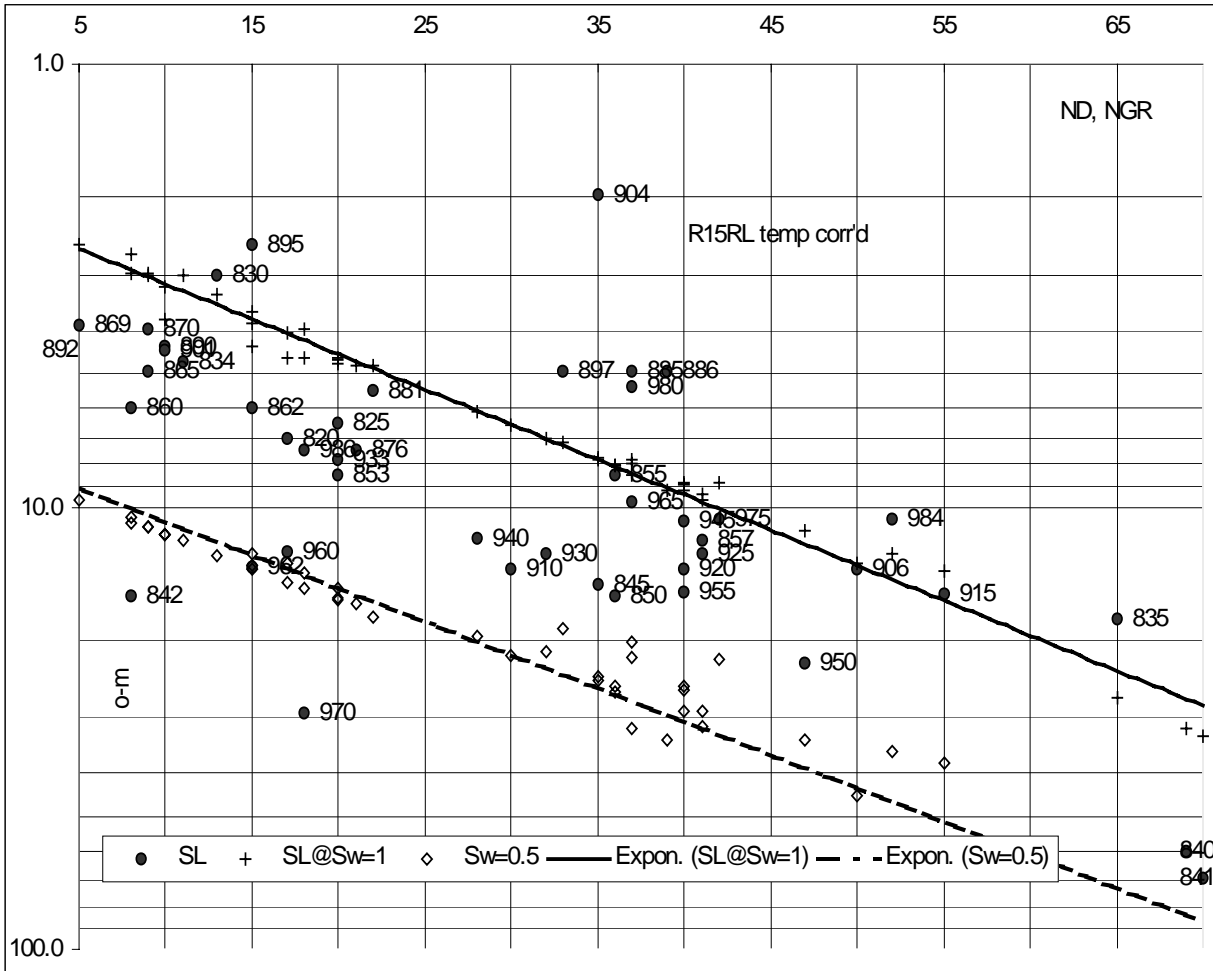


Normal Electrical Survey      Figure 8  
 Plot of Normal Electrical Survey Tool, Rn to Determine Rs and Rmf



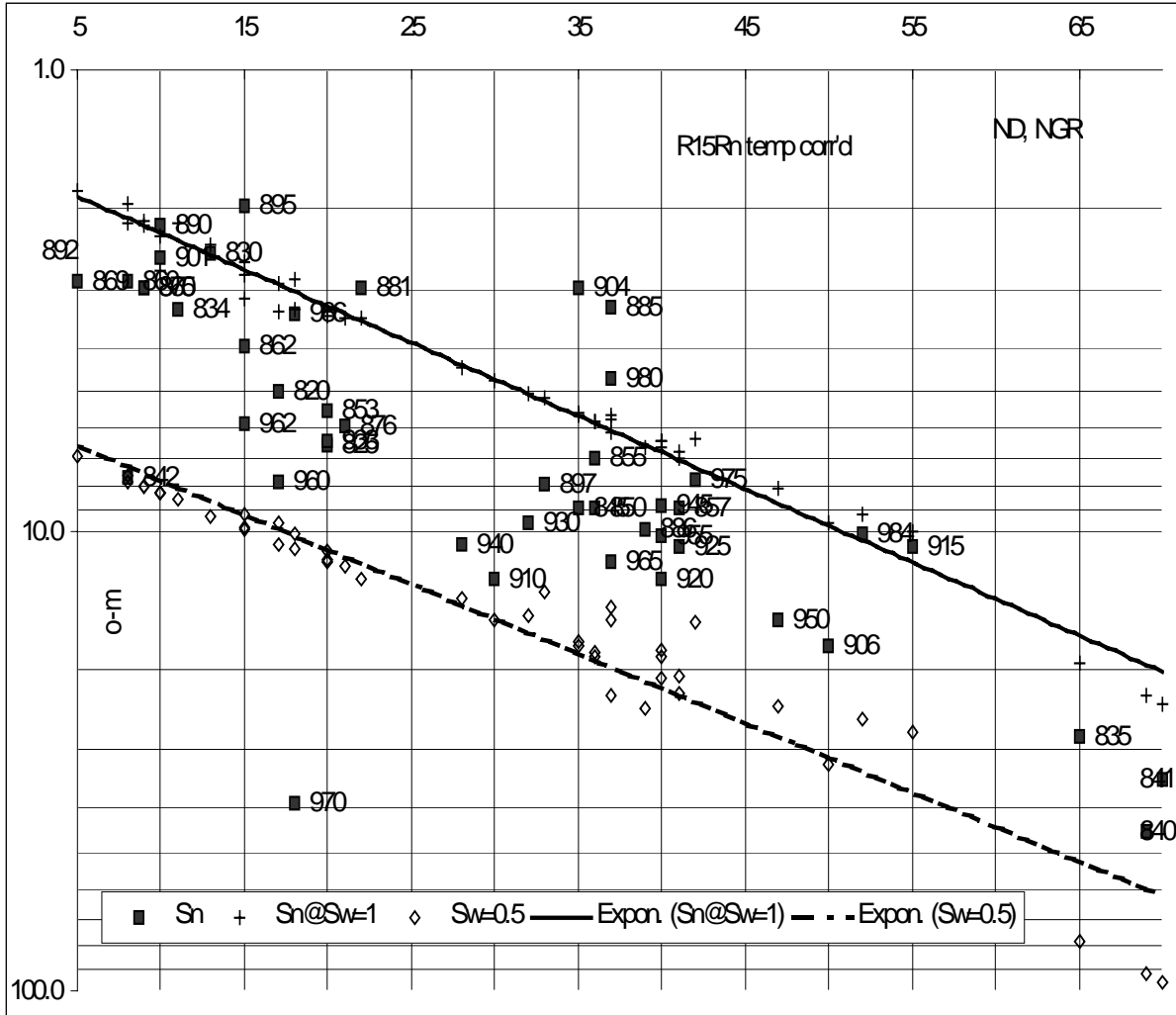
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Figure 9  
 Well K15 Pickett Plot for RL with Shale Correction Points Ordv Section



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Figure 10  
 Well K15 Pickett Plot for Rn with Shale Correction Points  
 Ordv. Section



APPENDIX

Appendix Background

Application of shale conductivity equations, permitted resolution of a substantial contradiction in water mineralization calculated from core open porosity, and particular attention is presented about this shale method. A good amount of technical literature presents rules for when to not use shale conductivity analysis, perhaps due to complications arising in the basic analysis equations. In the course of review, it seems that that no-one-single-factor is presented to represent the degree of shale effect on conductivity in porous media. This discussion presents a simplified approach using a new variant to the classic Archie conductivity equation,  $(S_w)^2 = F \cdot R_w / R_t$ , namely, adding F sub shale,  $F_s$ .

$$(S_w)^2 = F_o \cdot F_s \cdot R_w / R_t, \quad F_s = \{5C^2\} / [1 + 2C(F_o R_w / S_w)(V/R)_s] \quad \text{Eqn. A13}$$

$$S_{w-xo}^2 = F_o \cdot F_{sxo} \cdot R_{mf} / R_{xo} \quad F_{s-xo} = \{5C^2\} / [1 + 2C(F_o R_{mf} / S_{xo})(V/R)_s] \quad \text{Eqn. A14}$$

This single factor is valuable because it allows a rapid estimation of the degree to which shale conductivity pathway affects water saturation values. As  $F_s$  approaches one, effects of shale conductivity are reduced on overall conductivity. Also, presented is a method to predict simultaneously, shale and water resistivity for Laterolog. Prediction of shale resistivity is a noted weakness in application of shale analysis, Ref.1.

What the  $F_s$  term does is to compensate for the increased calculated-resistivity from reduction of total porosity by the shale volume. Shale Analysis increases F by using open porosity, which is, total porosity, reduced by shale volume. For example, a shale section gives a large NGR porosity, although the open porosity is very low. Without the  $F_s$  term, the classic Archie equation using open porosity would calculate  $R_t$  as an erroneous large number.

Appendix Summary

Modification of the classic Archie conductivity equation by  $F_s$  introduces a new canon of analytical equations. Wylie (p57) initially proposed a similar concept for analysis of shaly sections with SP data by “apparent water resistivity”. Application of  $F_s$  introduces two new “apparent conductive fluids” namely, “apparent water” and “apparent mud filtrate”

$$R_w\text{-a} = R_w \cdot F_s, \quad \& \quad R_{mf}\text{-a} = R_{mf} \cdot F_{s-xo}, \quad \text{Eqn A13 \& 14,}$$

When either  $F_s$  or the ratio  $F_s/F_{s-xo}$  deviate significantly from unity, then analytical equations need to use the “apparent water resistivity”,  $R_w\text{-a}$  and/or the “apparent mud filtrate resistivity” These analytical equations are variants of the two term shale conductivity equation, Simandoux modification of Wylie’s three term shale

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conductivity equation. The restrictions for dropping this third term should be carefully reviewed, together with the validity of the shale terms.

The following two term conductivity equation allows a fast evaluation with the basic constraint that L.H.S. term must be greater than or equal to zero:

$\{5C^2/(Rt) - (2CSw)(V/R)s\}$	$= (Sw)^2/(FRw)$	& is > or =0	A14b
$\{5C^2/(Rxo) - (2CSwxo)(V/R)s\}$	$= (Swxo)^2/(FRmf)$	& is > or =0	A14c

The use of A14b is more complicated than Eqn. A14c. This is due to the increased investigation depth of the Rt tools. This means increased chances that geometric factors (thin beds) or tool will encounter low Sw and, Sw is far less predictable than is Sxo, due to mud filtrate flushing. Sxo will be either 1 or if HC's present, then rarely below 0.5 and typically in the range of 0.85 to 1.0, Asquith pg 45. If after allowing for HC saturation, the LHS is consistently negative, a systematic error in either Vs or Rs is to be suspected.

A re-arrangement of A14c, allows for calculation of open porosity by:

$Rmf*\{5C^2/(Rxo) - (2CSwxo)(V/R)s\}/(Swxo)^2$	$= 1/(Fo)$	A14c
--	------------	------

Such open porosity values should roughly correspond to porosity tool responses, corrected for shale effects. Asquith, p103, points out that density porosity is the least susceptible to shale, provided shale density is close to matrix density. It can be stated that at Vs =1, open porosity is nearly zero, A14c, thus:

tool porosity =  $\phi_t = \phi_o + (V\phi_t)s$  &  $\ln\phi = aND+b$  A14d-1 &-2

In this instance, porosity, when available, was recorded in deflection units by a Neutron Gamma Ray , NGR, tool. The two point calibration method was used, with a shale porosity of 33%. Eqn A14c determined open porosity range and A14d gave max and min tool porosity response. Finally, neutron deflection, ND and shale volume at each depth point is used to calculate open porosity for all depth points. The porosity, shale volume, and temperature corrected Rs, Rmf or Rw at each depth is used to calculate the formation Sw =1 & Sw = 1/2 trend lines (Eqn A34 to A36) for the Pickett Plots. An example is given in Figure 9 and 10

$Rt = F_s * F * Rw / Sw^2$	$Rxo = F_{sxo} * F * Rmf / Sw^2$	A34a,b
$F = (1/ \phi^2) \text{carbonates } C= 0.45$	$(0.81/ \phi^2) \text{consol.sands } C=0.40$	A35a,b
$F_s = 5C^2 / (1 + 2C * (F * Rw / Sw) * [V/R]s)$	$F_{sxo} = 5C^2 / (1 + 2C * (F * Rmf / Sw) * [V/R]s)$	A36



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A.I.1: Derivation of F sub shale, Fs, Equation from Simandoux Equation

$$S_w = C * R_w * (1 / \Phi^2) [\sqrt{\{5(\Phi^2)/(R_w R_t) + (V_s/R_s)^2\}} - (V_s/R_s)] \quad \text{Eq.A1}$$

$$(\Phi^2) S_w / (C * R_w) = [\sqrt{\{5(\Phi^2)/(R_w R_t) + (V_s/R_s)^2\}} - (V_s/R_s)] \quad \text{A2}$$

$$(\Phi^2) S_w / (C * R_w) + (V_s/R_s) = [\sqrt{\{5(\Phi^2)/(R_w R_t) + (V_s/R_s)^2\}}] \quad \text{A3}$$

Square terms to remove radical

$$(\Phi^4) (S_w / (C R_w))^2 + 2(\Phi^2) (S_w / (C R_w)) * (V/R)_s + (V/R)_s^2 = \{5(\Phi^2)/(R_w R_t) + (V_s/R_s)^2\} \quad \text{A4}$$

the term  $(V_s/R_s)^2$  may now be subtracted from both sides leaving:

$$(\Phi^4) (S_w / (C R_w))^2 + 2(\Phi^2) (S_w / (C R_w)) * (V/R)_s = \{5(\Phi^2)/(R_w R_t)\} \quad \text{A5}$$

Multiply both sides by  $1/(\Phi^4)$

$$(S_w / (C R_w))^2 + 2(S_w / ((\Phi^2) C R_w)) * (V/R)_s = \{5/((\Phi^2) R_w R_t)\} \quad \text{A6}$$

Multiply both sides by  $R_w^2$

$$(S_w / (C))^2 + 2(R_w * S_w / ((\Phi^2) C)) * (V/R)_s = \{5 R_w / ((\Phi^2) R_t)\} \quad \text{A7}$$

Multiply both sides by  $C^2$

$$(S_w)^2 + 2(C R_w * S_w / ((\Phi^2))) * (V/R)_s = \{5 C^2 R_w / ((\Phi^2) R_t)\} \quad \text{A8}$$

The term  $(1/\Phi^2)$  may be recognized as F, the Archie formation factor

$$(S_w)^2 + 2(C F R_w * S_w) * (V/R)_s = \{F R_w / (R_t)\} * 5 C^2 \quad \text{A9}$$

$$\text{Aside: } 1/R_t = (2/5C) * S_w * (V/R)_s + 1/\{5C^2\} * 1/F R_w I \quad \text{Eq.IV \& A10}$$

On the LHS, bring out and isolate  $(S_w)^2$  to get the classic Archie Form

$$S_w^2 (1 + 2C(F R_w / S_w) * (V/R)_s) = \{F R_w / (R_t)\} * 5 C^2 \quad \text{A11}$$

$$S_w^2 = \{F R_w / (R_t)\} * \{5 C^2\} / [1 + 2C(F R_w / S_w) * (V/R)_s] \quad \text{A12}$$

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This is seen as the classic Sw Eqn with, **F<sub>s</sub>**, called F sub shale.

$$\mathbf{F_s} = \{5C^2\}/[1 + 2C(FR_w/S_w)*(V/R)_s] \quad \text{A13}$$

$$S_w^2 = \mathbf{F_s} * F * R_w / (R_t) \quad \text{A14}$$

An important secondary equation is the conductivity equation, arrived at from A9, starting with dividing both sides by (FR<sub>w</sub>) to get:

$$\{5C^2/(R_t) - (2CS_w)(V/R)_s\} = (S_w)^2/(FR_w) \quad \text{A14b}$$

Equation 14b appears significant in two ways: First, the R.H.S. must always be positive, so the LHS must always be greater than or equal to zero. Thus evaluation of the LHS for sections where Sw=1, should give a check on validity of shale volumes and shale resistivity value. Second, when F becomes large and at low Sw levels, the RHS will approach zero. Which seems to form basis of so called “dual water, (1/Sw = 1/Sw-s + 1/Sw-a)” models. For example calculation of Sw using the LHS leads to:

$$S_w-s = 2.5C(R/V)_s(1/R_t) \quad \text{A14d}$$

Compare A14d to the empirical Sw-s term of dual water Indonesian model:

$$S_w-s = \{V_s^{(V_s/2-1)}\} * (R_s/R_t)^{1/2} \ \& \ V_s = (R_s/R_t)^\alpha \ \alpha=1 \ \text{or} \ 2 - (2R_s/R_t)^{0.25} \ \text{if} \ R_s/R_t < 0.5$$

The third conclusion from A14b, is that as V<sub>s</sub> approaches 1, the RHS must approach zero for the open porosity goes to zero. In this case at Sw=1, then

$$R_s = 0.4/C * (R_t)(V_s) \quad \text{Eqn A14C}$$

The results of A14C is that many instances, R<sub>s</sub>, is taken simply as resistivity of an adjacent shale bed. This method is exceptionally simple, provided there are no geometrical effects such as washed out bore hole or thin beds, and sectional resistivity represents the type clay intermingled with the formation of interest. Asquith notes that montmorillonite and illite lower resistivity more than kalonite and chlorite shale.

The key element of equation A13 and 14, is that permutations of R<sub>t</sub> or R<sub>xo</sub> may be rapidly calculated for Pickett plots. The method applied here is: given, ND; Neutron log Deflection, V<sub>s</sub>, R<sub>s</sub>, and R<sub>w</sub>, then calculate R<sub>t</sub> for Sw=1 and Sw= ½ these 2 points are then plotted at the total porosity, ND point, along with the deep resistivity value. A similar second plot of R<sub>n</sub> using R<sub>mf</sub> is produced to look for hydrocarbon mobility or for thin beds where the full value of deep resistivity would not develop. An example is given in Figures 9 and 10. The section from 905 to 980 m of Fig's 9 and 10 bailed at approximately 1 ton oil per day and exhibits an oil static top pressure of about 40 psi above grade.

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A.I.2: Determination of Rs & Rw When F is Known

The conductivity equation for Laterolog may be reduced at Sw=1 to:

$$R_w/R_t (5FC^2) = 1 + 2CV_s(FR_w/R_s) \text{ or} \quad \text{Eq.A15}$$

$$F/R_t = FV_s 2/(5CR_s) + 1/(R_w 5C^2) \quad \text{Eq.A16}$$

When data is plotted on these x, y coordinates as y=mx+b, then slope m and intercept b will calculate Rw and Rs.

$$R_w = 1/(5bC^2) \quad \text{and} \quad R_s = 2/(m5C) \quad \text{Eq.A17-a,b}$$

A.I.3. Determination of F from Shale Factors when Rw & Rs are known

Consider the case of a porous solid mixed and consolidated with modest amounts of clay particles . If one measures the porosity of this solid by resistivity in a saline solution, the calculated porosity of the new solid would be increased, directly in proportion to the fraction of consolidated clay introduced. "In all cases... the effect (of clay or shale) is similar. The clay acts as a separate conductive path additional to that afforded by the saline solution in the rock pores" , Wylie pg.16. This effect was quantified by Wylie with the following conductivity equation:

$$1/R_t = A \cdot S_w + 1/(F \cdot R_w \cdot I) + B \quad \text{Eq-1}$$

Equating Wylie's conductivity equation (for Laterolog) to Simandoux equation, Eqn.A10, one arrives at C=0.45 and the following conductivity equation, when Sw=1 :

$$1/R_t \approx 1/(FR_w) + 0.9(V/R)_s \quad \text{Eq.A18}$$

$$F/R_t = 1/R_w + 0.9(F \cdot V_s)/R_s \quad \& \quad R_s=0.9/m \quad \text{Eq.A19a,b}$$

$$1/F = R_w/R_t - 0.90 \cdot R_w \cdot (V/R)_s \quad \text{A20}$$

or for Rn and Rmf

$$1/F = R_{mf}/R_n - 0.90 R_{mf} \cdot (V/R)_s \quad \text{A21}$$

A.I.4 Sectional Summary

In summary, this appendix proposed an analysis of well log data by the Pickett Plot, (PP) while allowing for considerations about shale levels. The Picket Plot looks for anomalies in resistivity by plotting log of resistivity against Neutron tool deflections, ND. Once anomalies are noted on the PP, it is useful to have an estimate of the degree of water saturation. In most instances, the presence of shale reduces resistivity deflections. This in turn makes hydrocarbon detection more difficult on a PP. It was proposed to add 100% & 50% Sw points to the Picket plot using a re-arranged Simandoux equation.

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APPENDIX II

What follows are considerations on the parametric factors and paradigms related to shale section analysis. The result is to offer a more comprehensive parameter for application of shale analysis. Finally, it contrasts the overly cautious position taken by Asquith: “ if a geologist overestimates shale content, a water bearing zone may calculate like a hydrocarbon zone”. In contrast, if shale factors are ignored, potential hydrocarbon bearing sections may be overlooked. In zones of low shale factor the Simandoux equations calculate exactly like the classic water saturation equations and thus, there is no loss of functionality when applying the proposed method, as contrasted to an overly cautious approach.

A.II.1 Conductivity Equation Review

Wylie proposed the following conductivity model of porous media with shale acting as a parallel conductor to the water inside the pores.

$$1/R_t = A \cdot S_w + 1/(F \cdot R_w \cdot I) + B \quad \text{Eq-I}$$

Wylie’s term B, represents conductivity thru paths involving clay only. Wylie states that “most often the clay is dispersed and B can be ignored, but that when there are thin or continuous clay streaks, B can be quite large”. The middle term is the standard conductivity term of porous media filled with conducting and non conducting phases. The first term is conductivity of brine filled pores and dispersed clay particles in series. Implicit in this equation is: “shale conductivity acts independent of the immersing water salinity”.. The middle term is the normal conductivity term when shale is ignored.

Simandoux refined this concept into a widely accepted equation:

$$S_w = C \cdot R_w \cdot (1/\phi^2) [\sqrt{\{5(\phi^2)/(R_w R_t) + (V_s/R_s)^2\}} - (V_s/R_s)] \quad \text{Eq-II}$$

It was shown that Simandoux equation can be arranged as:

$$1/R_t = (2/5C) \cdot S_w \cdot (V/R)_s + 1/\{5C^2\} \cdot 1/\{F \cdot R_w \cdot I\} \quad \text{Eq-III}$$

In this form, it is clear Wylie’s B term was dropped. To the extent this equation is a close model of actual physical process for conductivity in porous media., it is possible to re-arrange the terms to analyze well log data, especially on the Picket plot as shown above.

AII.2 Comparisons of Wylie Core Data to Simandoux Equations

Wylie and the AAPG manual conclude that shale effects are marginal unless salinity levels exceed 0.3 o-m at 80F. For example, when plotting Wylie’s core analysis data for Burgan Sands, against Simandoux Equation with  $F \sim 5$  and between  $V/R=0.6$  and

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0.19 respectively, Figure 3, one could conclude that at low  $R_w$  values, shale has little effect on calculated  $S_w$  values. As is shown in Figure 3, for  $R_w$  less than 0.3, both Wylie's data and Simandoux equation have a resistivity index close to 1.

However, this is only true at small values of  $F$ . For low porosity formations this is not so. Possibly because the shale conductivity path is a greater % of the total conductive path. Notice when  $S_w=1$  that  $F$ : observed, is  $R_t/R_w$ , then equation 1 is:

$$F_a/F_o = 1 + 0.9(F_a)(R_w)(V/R)_s \quad \text{Eq.IV}$$

When  $F_a/F_o$  vs  $F$  is plotted, Figure 4, it can be seen that at large Formation Factors (right hand side of Fig4), as with many carbonate systems, substantial errors can result, if shale corrections are ignored, even when water resistivity ranges 0.06 to 0.14, even with the modest shale values of Figure 4:

Wylie presented some core analysis in the form of  $I$  vs  $S_w$ . This appendix shows that  $I$  can be calculated in terms of the shale factors, in which instance  $I$  is called  $I_s$ ,  $I$  sub shale.

$$I_s = \{1/(S_w^2)\} / [1 + 2C(FR_w/S_w)*(V/R)_s] * [1 + 2C(FR_w)*(V/R)_s] \quad \text{A33}$$

Some important parametric points come from the  $I$  sub shale equation:

At  $R_w=0$  or nearly 0.00, then  $I_s = I = 1/(S_w^2)$ , for any  $F$  or  $(V/R)_s$

At  $S_w = 1$  the  $I_s = I = 1$  for any  $F$  or  $(V/R)_s$  or  $R_w$

As  $S_w$  approaches 0.00,  $I$  approaches  $1/S_w$  for any  $F$  or  $(V/R)_s$  or  $R_w$

As  $S_w$  decreases, the effective power of  $S_w$  will range between  $-2$  &  $-1$

Increases in either  $F$  or  $(V/R)_s$  can offset any parametric rules for  $R_w$

This last point is illustrated by plotting Wylie core studies and a parametric change in  $F*(V/R)_s$ , Figure 5.

Again, to match Wylie's core data, an  $F$  on the order of the Burgan Sands  $F$ , 5, is used. Figure 5 shows that with low  $F$  values studied by Wylie, there is not a need to correct for shale effects when water resistivity is less than 0.30. However, with modest  $F$  values, and  $(V/R)_s$ , shale correction factors become more important. It is seen in Figure 5 at  $I = 3$ , an  $S_w$  of 60% is actually 50% and at  $I = 10$ , the un-corrected  $S_w$  is 32% vs the corrected value of 19%.

Another paradigm proposed by Asquith, was "that for shale content to significantly affect log derived water saturations, shale content must be greater than 10 to 15%".

Figure 5 shows this statement to be inaccurate when open porosity falls below about 15%. For example, at 5% open porosity, ignoring shale effect, gives  $S_w = 1$ , and at 10% open porosity, ignoring shale factors, one arrives at 30%  $S_w$ , based on a modest shale factor of between 9 and 13%. The correct parameter to define cut off points is  $F_s$ , not simply  $V_s$ .

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A.II.3 Formation Resistance For Picket Plots

Once formation water resistance,  $R_w$ , is known it is then possible to calculate the Formation Resistance,  $R_t$  based on the Simandoux variant of Wylie equation; given an assumption of formation lithology and porosity.

$$(S_w)^n = F' \cdot F_s \cdot (R_w/R_t) \text{ or } R_t = F' \cdot F_s \cdot (R_w/(S_w)^n); \quad \text{Eqn V}$$

$$\text{At } S_w = 1.00, \text{ then } (S_w)^n = 1.00 = F' \cdot (R_w/R_t) \text{ so} \quad \text{Eqn VI}$$

$$R_t = F' \cdot F_s \cdot R_w / (S_w)^2 \quad F_s \text{ relates to porosity \& lithology.} \quad \text{Eqn VII}$$

For Pickett plots the Eqn is arranged to calculate  $R_t$  at  $S_w=1$  and  $S_w = 1/2$  so as to allow judgment of the degree of water saturation, while accounting for shale and open porosity effects.

A.II.4 RESISTIVITY RATIO WITH SHALE CONSIDERATIONS

One key consideration in log analysis is that formation factors cancel out when considering resistivity ratios in the same matrix, but with different fluids. However, when shale are present,  $F_s$  cannot cancel as does the standard  $F$ , unless  $R_{mf} = R_w$ . As a rule, only small sections of a well ever have zero SP. Typically mud is formulated more saline than formation waters to reduce clay swelling tendency and thereby maintain formation porosity. When  $R_{mf}$  does not equal  $R_w$ , or when  $S_w$  for the two sections are not equal,  $F_s$  does not cancel and consideration should be given to the ratio of  $F_s$ . For example, at  $S_w=1$ ,

$$S_w=1 = F_s \cdot F \cdot R_w / (R_t) = F_s \cdot F \cdot R_{mf} / R_{xo} \quad \text{giving:} \quad \text{A22 a,b}$$

$$\{5C^2\} / [1 + 2C(FR_w/S_w) \cdot (V/R)_s] \cdot F \cdot R_w / (R_t) = \{5C^2\} / [1 + 2C(FR_{mf}/S_w) \cdot (V/R)_s] \cdot F \cdot R_{mf} / R_{xo}$$

Since lithology and porosity are considered uniform, radial to the well bore,  $C$  and  $F$  cancel, leaving:

$$\{1\} / [1 + 2C(FR_w/S_w) \cdot (V/R)_s] \cdot R_w / (R_t) = \{1\} / [1 + 2C(FR_{mf}/S_w) \cdot (V/R)_s] \cdot R_{mf} / R_{xo}$$

$$R_w / R_{mf} = (R_t / R_{xo}) \cdot [1 + 2C(FR_w) \cdot (V/R)_s] / [1 + 2C(FR_{mf}) \cdot (V/R)_s] \quad \text{A25}$$

Considering a carbonate situation with  $F=100$ ,  $R_w = 0.12$ ,  $R_{mf} = 0.1$ ,  $C=0.45$ , and  $(V/R)_s = 0.11$ , one arrives at

$$R_w / R_{mf} = (R_t / R_{xo}) \cdot [2.06] / [1.88] = 1.1 \cdot (R_t / R_{xo})$$

Considering a sandstone situation with  $F=35$ ,  $R_w = 0.06$ ,  $R_{mf} = 0.1$ ,  $C=0.4$ , and  $(V/R)_s = 0.05$ , one arrives at

$$R_w / R_{mf} = (R_t / R_{xo}) \cdot [1.084] / [1.14] = 0.95 \cdot (R_t / R_{xo})$$

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The Table below uses data from Table 1 to estimate Ordovician water resistivity. These calculations arrive at  $R_w$  about 25% less than if shale factors are ignored.

	Ord	Ord	Ord	Cmb	Ord	Ord
	RL	Rn	Rt/Rxo	RL	Rn	Rt/Rxo
avg	10.1	13.3	0.8	1.8	6.7	0.3
F's	2.05	3.24		1.07	1.30	
Rs			2.20			2.80
Vs			0.21			0.12
C			0.45			0.40
F			90.00			30.00
$R_w$			0.14			0.065
Rmf			0.29			0.29
V/Rs			0.10			0.04
<b>Rmfc</b>				<b>calc'd=</b>		<b>0.29</b>
<b>Rwc</b>		<b>calc'd=</b>	<b>0.14</b>			

**A.II.5 RESISTIVITY INDEX WITH SHALE CONSIDRATIONS**

The resistivity index, I, is an important concept for log analysis.

$$I = R_{@Sw} / \{R_{@Sw=1}\} = R_t / R_o = R_t / \{FR_w\} = kSw^{(-n)} \quad A27$$

Generally,  $k=1$  and  $n = 2$ , although an analysis of Wylie's' core analysis for San Andres, Texas carbonate cores gave

$$k=0.5 \text{ and } n = 2 \text{ where } F = 1.25/\phi^{2.2} . \quad A28$$

In the case where shale is involved, I can be defined as follows:

$$\text{Since the resistivity index, } I, = \{R_{@Sw}\} / \{R_{@Sw=1}\} \quad A29$$

Simandoux equation can be recast as:

$$R_t = \{5C^2\} \{FR_w / (Sw^2)\} / [1 + 2C(FR_w/Sw) * (V/R)s ] \quad A30$$

$$(R_{@Sw=1}) = \{5C^2\} \{FR_w / (1)\} / [1 + 2C(FR_w/1) * (V/R)s ] \quad A31$$

$$\{R_{@Sw}\} = \{5C^2\} \{FR_w / (Sw^2)\} / [1 + 2C(FR_w/Sw) * (V/R)s ] \quad A32$$

$$I_s = \{ \{1 / (Sw^2)\} / [1 + 2C(FR_w/Sw) * (V/R)s ] \} * [1 + 2C(FR_w) * (V/R)s ] \quad A33$$

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**A.II.6 PICKETT PLOTS**

The Pickett Plot, PP, is a plot of formation R on a log scale y axis and Neutron Deflection on a linear x scale. In order to evaluate the degree of hydrocarbon contained in the rocks, it is necessary to plot a family of water saturation lines also. When shale is not present, this is a straight forward process. When shale may be present, construction of the Sw lines on PP's can be accomplished by re-arranging the above to:

$$R_t = F_s * F * R_w / S_w^2 \quad A34$$

The two lines of interest on the PP are at Sw=1 and at Sw = ½ although additional lines are possible, these two points generally suffice to flag zones of interest for more detailed analysis. The term F is evaluated with the suitable form, given, lithology

The AAPG recommended equations for F are:

$$F = (1 / \phi^2) \quad \text{carbonates and } (0.81 / \phi^2) \text{ consolidated sands Eqn A35a,b}$$

The additional shale factor uses, F<sub>s</sub> Simandoux.

$$F_s = 5C^2 / (1 + 2C * (F * R_w / S_w) * [V/R]_s) \quad C = 0.45 \text{ Carbs, } 0.40 \text{ sands. Eq. A36}$$

When using shale analysis, F must be calculated on shale free, open, porosity basis, Equations for porosity corrections and shale volume are given in AAPG manual. Shale volume is calculated from gamma ray index, and R<sub>s</sub> is resistivity of 100% shale bed, see Eqn A15/16.

It was found that a better trend (R<sub>t</sub> at Sw= 1 and ½ ) develops by constructing two separate plots of Ordovician and Cambrium ND and RL.

If an API calibrated ND tool is used, then calculation of limestone total porosity units is direct:

$$\text{Log}(p, v/v) = 0.73 - 0.00143 \text{ND}_{\text{api}} \quad A37$$

And in the general form, when calibration is not available,

$$\text{Log}(p, v/v) = U - V * \text{ND} \quad A38$$

In analysis of these logs, no such calibration existed, and it was necessary to use the general form, A38. The porosity units were calculated for sections with good caliber by taking ND minimum and maximum corresponding to about 26% and 5% total porosity respectively. For Cambrium sandstone sections, 4% was added to the calculated porosity units.

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The basic equations for PP analysis w/o shale factors are as follows:  
Given a regression line for SW =1 of the form:

$$\text{Log}(R_t) = M \cdot ND + B \quad \text{A39}$$

$$\text{At } F = 10, R_w = R_t/10, \text{ and } R_{mf} = R_{xo}/10 \quad \text{A40a, b}$$

use Eq.A35a or b to calculate porosity for F=10, then apply A38 to get ND, and lastly, apply 40a,b to arrive at R<sub>w</sub> or R<sub>mf</sub>.

When this method was applied to Ordovician section of 7 well logs, the average R<sub>w</sub> was 0.13, with a range of 0.09 to 0.18 o-m. This method made an educated assumption about the porosity range.

If R<sub>w</sub> and R<sub>s</sub> are known, as postulated by this paper, then either, the basic Archie or the F<sub>s</sub> modification can be used to determine total porosity range and apply A38 to scale open porosity units.

Once open porosity units are scaled to ND, then shale corrected R at Sw=1 and ½ points are plotted at each original ND along with log resistivity. For example see Figures 9 and 10.

### APPENDIX III Discussion of Porosity Factors

Simply, there are two types of physical pores: 1) open pores which communicate flow (V<sub>o</sub>) and 2) closed pores which do not communicate flow (V<sub>c</sub>). Those which do not communicate flow can be by virtue of closure (poorly connected Vugs), by virtue of size (Micro Pores), and capillary action. Differentiation of these two pore types is loosely made by the terms “total porosity, φ<sub>t</sub>” and “open porosity φ<sub>o</sub>”. The total porosity being composed of the sum of the two terms V<sub>o</sub> + V<sub>c</sub>, which can be said to have a sum of 1, with a total matrix volume (V<sub>t</sub>) being sum of solids, open holes and closed holes. Porosity is the fraction of pores to total volume:

$$\phi_t = (1)/V_t \ \& \ \phi_o = V_o/V_t, \ \text{so } \phi_t/1 = \phi_o/V_o = 1/V_t \ \& \ V_o = (1 - V_c) \ \text{so } \phi_o = \phi_t * (1 - V_c)$$

In the case of a solid having a uniform pore fabric, with no vugs or micro pores, then open porosity is equal to total porosity.

Sometimes, open porosity is taken as φ<sub>o</sub> = φ<sub>t</sub>\*(1-V<sub>c</sub>) or φ<sub>o</sub> = φ<sub>t</sub>\*(1-V<sub>sh</sub>). Where V<sub>sh</sub> is volume of shale introduced, shale being a consolidated clay. This last equation is Asquith’s simple approach to correct φ<sub>n-d</sub> for shale, pg 188, with results close to that obtained by more detailed methods.

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Generally, porosity correction formula take the form:  $\phi_o = \phi_t - \phi_{ts}V_s$ . This forces the value of open porosity to zero at  $V_s = 1$ , because tool response,  $\phi_t$ , is  $\phi_{ts}$  at  $V_s = 1$ . This correction is necessary because tool porosity response is calculated without regard to shale. Sometimes, shale density (2.4 to 2.8 g/cc) will be about equal to matrix density. In this case, no correction is required to the density porosity response. To a lesser extent, this may hold for sonic porosity, in the case of consolidated shale. Neutron tools respond to hydrogen content and the accepted calibration point for shale ranges between 30 (Schlumberger with 45% porosity adjacent clay, Asquith pg 103) and 33% (Wylie, pg 120).

Wylie states that in sedimentary rocks there is virtually no non-effective porosity, i.e. all pores contain water directly interconnected with water in other pores.” However from a producing view point effective porosity is that quantity of pores which can contribute to fluid flow. Water contained by shale cannot contribute to flow due to the very low permeability factor. Likewise “connate water” cannot contribute to flow. Connate water in well log literature is often called  $S_w\text{-irr}$ . For flow calculations connate water is considered as rock, effective porosity would possibly be  $\phi_e = (\phi_t - \phi_{ts}V_s) * (1 - S_w\text{-irr})$ . Hamada, Ref 15, considers there to be a lack of methodology in electrical logs to differentiate between shale waters and waters on pore walls held by static forces. Pirson<sup>17</sup>, presented a modified Wylie method (using electrical tools only) and calculates shale hydrated water as:

$$V_{ws} = (R_{mf}/R_w)^{V_{sh}} - 1 / (R_{mf}/R_w - 1) \quad \& \text{ calculate permeabilty by tortuosity method.}$$

For Pirsons' example at  $R_{mf}=0.42$ ,  $R_w=0.06$   $V_{sh}=0.71$  and  $S_w=0.542$ , then  $V_{ws} = 0.497$ , the free water is  $0.542 - 0.497 = 0.045$  and the well completed with water free production.

**Summary:**

The proposed term,  $F_s$ ; Eq. A36 is useful in defining the degree of impact shale factors can have on water saturation calculations. The apparent water ( $R_w * F_s$ ) or apparent mud filtrate resistivity ( $R_{mf} * F_{sxo}$ ) are useful diagnostic parameters when evaluating validity of ratio methods in shale sections.

### Discussion of First Generation e-Log Tools

Most Latvia e-log data are from a Normal (2 electrode) sonde N2M0.25A, plus a Lateral (3 electrode) sonde A2M0.25N. In some instances a 0.5m spacing was used for a Normal log. Characteristics of these first generation devices are seldom covered in modern geology. Modern logging tool has tools to read  $R_{xo}$  and  $R_t$  with a minimum of correction.

There are 2 types of correction factors, environmental and geometric. Environmental factors are those factors relating to resistivity of: mud  $R_m$ , formation  $R_i$ , invasion zone  $R_i$ , flushed zone  $R_{xo}$ , and shoulder  $R_s$ . Geometric factors are those related to hole diameter, sonde diameter, bed thickness, and invasion diameter.

In both Normal and Lateral instruments, points A and B are current electrodes and M and N are voltage electrodes. For lateral Sonde  $AO=(AM+AN)/2$  and investigation depth is about  $AO$ . Investigation depth is  $2AM$  for the Normal sonde. Both sondes read  $V=(R_i/4*3.14)(1/AM-1/AN)$ . In a USA Normal arrangement  $1/AN$  is insignificant compared to  $1/AM$  and  $V$  is taken as just  $(R_i/4*3.14)(1/AM)$ . For LV normal sonde this is not the case.

A feature of an unfocused sonde is current saturation. For lateral tool borehole correction is not significant for  $R_a/R_m < 50$ , and  $R_a/R_m = R_t/R_m$  for all but extreme values of  $(s/d_h)$ , sonde spacing/hole diameter. Likewise for the normal sonde borehole correction is not significant for  $R_a/R_m < 10$ . But in front of a highly resistive beds or when the ratio of  $(s/d_h)$  to is small,  $R_t/R_m$  will always be greater than  $R_a/R_m$ . Current saturation appears as the asymptotic value of  $R_a/R_m$  or  $R_a/R_s$  on departure charts, given a large value of ordinate value,  $R_t/R_m$  or  $R_t/R_s$ . Readings close to the asymptotic value should be considered unreliable..

For a 4 electrode lateral sonde, current is split:  $R_{max}/R_{mud} = (8*(AO/d_h)^2)/(1-(d_s/d_h)^2)$ . For an 8.75" hole and a 75mm sonde of  $AO=32"$ ,  $126 = R_{max}/R_{mud}$ , Fig.9.3, Pirson. For a 3 electrode lateral sonde current is not split and  $R_{max}/R_{mud} = (4*(AO/d_h)^2)/(1-(d_s/d_h)^2)$ . For an 8.0" hole and a 3.5" sonde of  $AO=18'8"$ ,  $825 = R_{max}/R_{mud}$ , Fig.2.15, p16, Hilchie. The formula for Normal sonde is  $R_{max}/R_{mud} = (8*(AM/d_h)(AN/d_h))/(1-(d_s/d_h)^2)$  and for Schlumberger sonde of  $AM=16"$ ,  $AN=240"$  and a 3.5" sonde in a 16" hole  $R_{max}/R_{mud}=125$ , Fig 3.8, pg.24 Hilchie. This relationship shows why longer electrode spacing are less effected by hole diameter  $d_h$ . These relationships allow determination of mud resistivity, given  $R_{max}$ ,  $d_h$ , and  $d_s$  or given  $R_{mud}$ , sonde spacing and diameter, determine hole diameter, below table. A short coming of the LV short normal arrangement is the relatively low saturation value, about 1/5 that of the US configuration. Corrections beyond 75% of the saturation value are too large to provide a reliable value. The below table list saturation values for the LV sondes in 8, 10, 16 inch holes.

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Changes in formation resistivity make the relationship between investigation radius and radius of invaded zone or flushed zone significant. Invasion

	Normal	Lateral		Well Rn max	Rmud 8"hole
sonde	N2M0.25A	A2M0.25N		Crk265/100	0.75
AO	na	2.125m	(7ft)	K15/44	0.33
AM us=	11.3	na	inch	Ed60/70	0.5
R/Rm max	136	545	8" hole 85mm ds	Ed17/50	0.38
R/Rm max	80	322	10"hole 85mm ds	K11&D13/40	0.29
R/Rm max	29	116	16"hole 85mm ds	K1/60	0.44
r.Invest'gn	0.50	2.00	Meter radius	K2/70	0.52

depth may be estimated as  $D_i/d_h = 25/\exp(\%p/8)$ . At porosity greater than 2%, the 2m lateral view of pore salinity will be that of natural waters.. At porosity less than 16%, the 0.25m Normal sonde, view of pore salinity will be that of mud filtrate, as detailed in below table.

The reading expected for Rn should be greater than  $R_L$  in the porosity range between 2% and 16% and for  $R_{mf} > R_w$ . At porosity under 2% both sondes should read invaded zone resistivity. Both sondes should read natural water salinity at porosity over 16% or if zero permeability.

%p/dHole"	4/8"	11/8	2/8	14/8
Invasion R	1.5m	0.64m	2.0m	0.44m

The invaded zone lies between the flushed zone and the unaffected zone. The salinity equation of the invaded zone is  $1/R_z = z/R_w + (1-z)/R_{mf}$ . And  $z = (\phi_v/v/2 - .01)$ . For low porosity systems and for  $R_{mf}/R_w < 10$ ,  $R_z = R_{mf}$  and  $R_i$  approximates  $R_{xo}$ .

One geometric factor is the formation thickness such that  $R_a = R_t$ . In the case of a Normal sonde, a uniform bed thickness equal or greater,  $e(m) = 12AM^{0.5} = 6m$  will read conclusively, without correction charts. The error is less than 15% for beds  $e/AM > 8$ . While thickness between AM and 2AM have unreliable resistivity values.

For Lateral sondes, beds greater than 2 AO read  $R_t$  by the midpoint method or , equal distant between the two main breaks in the curve; for  $e = 1.5AO$ , use 2/3 rule, for  $e = 1.3AO$  read  $R_{max}$ , for  $R < 1.05AO$  and  $> 0.85$  no reading is possible, for  $e > 0.25AO$  and  $< 0.5AO$ ,  $R_t > (R_s/R_{min})R_{max}$ . The lateral values so calculated should be adjusted for borehole and shoulder effects if  $R_{sn}/R_{mud} > 50$ . Other criteria are  $R_{mud} < 5R_s$  &  $> 0.3R_s$ ,  $d_{hole} > 0.01AO$  &  $< 0.1AO$ , mud invasion  $< 0.2AO$ , for beds of  $e < 1.5AO$  &  $> 1AO$  a determination is not plausible.

The 'N2M0.25A' normal sonde of Latvian logs, has spacing of 0.25m (10"). Normal sondes read a thick resistive bed short by  $AM/2$  on both top and bottom and will read a conductive bed long by  $AM/2$  on top and bottom, and in both instance have symmetrical curves. For a bed shorter than AM, it produces either 2 minimums or 2 maximums of length  $e + AM$ . For Guyod and  $e/AM = 3.6$ ,  $R_a/R_s = 7.8$ ,  $R_t/R_a = 1.6$  chart, 1.65, this equation. Guyod chart also indicates that beds greater than 2 m will read directly with errors less than 20%.

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Sonde	CF type	Equation	
Nlong	shoulder	$R_t/R_a = [(3.149/(e/AM)^{1.8}) \ln(R_a/R_s) + 1] e/AM > 2$	Guy. Fig3p139
N		$R_i/R_s = \exp(3.28 + .03R_a/R_s - .08e/AM)$	Pirson
Nshort	bore	$R_i/R_m = a(R_a/R_m)^m \quad m = 1.28/(AM/d_h)^{0.23}$ $a = 0.77 + 0.036(AM/d_h)$	Sch 1-3AM/d <sub>h</sub>
Lat		$R_t/R_a = (a)e^{(bR_a/R_s)} \quad a = 1.18 - 0.376e^{(b)}$ , $b = 0.83(e/AO)^{0.56}$ . valid: $0.1 > e/AO < 1$	Guy. Fig6p140
Lat	shoulder	$S = 3.45 - 1.53 \ln T$ ; $T = e/AO$ & $A = [(R_a/R_s)/S - 1]$	Guy. Fig.6.19
		$R_t/R_s = 35 \exp(A(1.2T + 4))$ for $e/AO < 1$	Guy. Fig.6.19

Pirson's interpretation rules, for the normal sonde are as follows:

Schlumberger give a chart for the short normal sonde regressed as follows:

$F_a = (R_n/R_m)^{1.22} (AM/d_h)^{-0.79} (R_{mf}/R_w)^{0.33}$  or  $\log F_a = 1.22 \log(R_n/R_m) - 0.79 \log(AM/d) - .0042SP$ , where  $SP = -75 \log(R_{mf}/R_w)$ , for  $R_n/R_{mud} = 7.9$ ,  $AM/hole, d = 2$  and  $SP = -108mv$ ,  $F_a = 20.2$ , if zone is thought to be HC flushed use  $F = F_a/S_i^2$ , ie at 30% residual HC  $S_i = 0.7$  &  $F = 40.4$ . For 30API oil  $ROS = \%POR$ , double this for gas or low API and at 45API the values may be halved..

Interpretation of logs run with a short normal depends on the invaded zone,  $S_i = \sqrt{(FR_z/R_i)}$ .

In the instance only  $R_{sn}$  is reliable, Pirson proposed  $S_w = F/(aR_i/R_{mf})$ . Where  $a = z(R_{mf}/R_w - 1) + 1$  and  $R_i$  is  $R_{sn}$  corrected for borehole and shoulder effects. It is based on Tixler  $S_{x0}$  eqn for  $S_w$  in the rocky mountain area, or  $S_{x0}$  equals root of  $S_w$ .

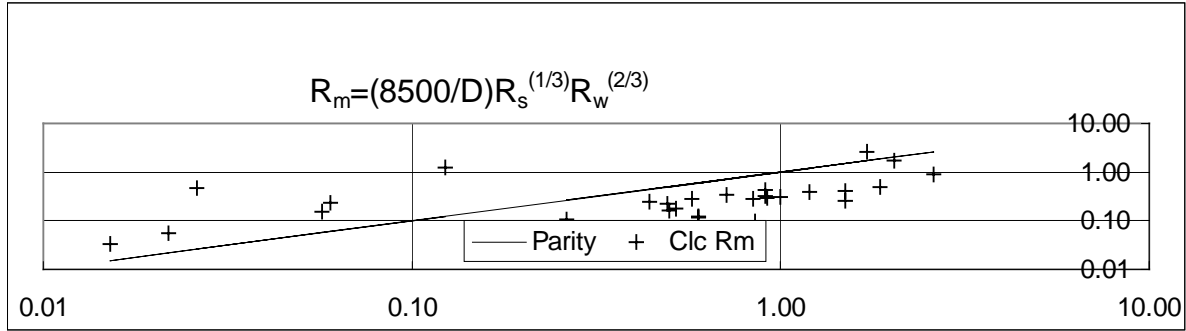
The lateral sond will read a thick bed short by  $AO$  and also produce a shadow zone of  $AO$  length below bed top., and for  $e < AO$  will also produce a reflection peak below the bed at distance  $e + AO$ .

A correction chart of Guyoid, Fig6.19 for Lateral sonde of  $e/AO < 1$  for resistive beds is as follows:  $T = e/AO$ , For  $R_a/R_s = 2.9$ ,  $T = 0.75$ ,  $R_t/R_s = 10.1$  vs 8.4 by chart. Reported valid for  $R_m > 0.3R_a$  and  $< 5R_a$ ;  $d_h > AO/100$  and  $< AO/5$ ;  $MN > AO/100$  and  $< AO/5$ .

Another Guyoid Lateral chart, Fig6p140, valid for  $e/AO < 1$  &  $> 0.1$ . for  $e/AO = 0.45$ ,  $R_a/R_s = 4.14$ ,  $R_t/R_a = 4.2$  chart, and 4.65 this equation.

The chart below is an empirical fit,  $R_m = (8500/D) R_s^{(1/3)} R_w^{(2/3)}$  of mud resistivity for a variety of US wells, using predominately natural muds. It was inspired by Guyoids' comment that  $R_s/R_m$  seldom exceeds 3. In this data set the median was 3, but ranged from 77 to 0.4, Depths 14400ft to 2500ft,  $R_w$  0.3 to 0.01,  $R_s$  7.0 to 0.40. In modern practice, the resistivity value of natural muds are modified to improve electric log accuracy. For Laterolog it is ideal for  $R_{mf}$  to equal  $R_w$  and for induction logging it is preferred that  $R_m$  exceed  $3R_w$ .

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In western Latvia, a comparison was made of Sw by Archie Method vs Sw by uncorrected ratio method. The results of the uncorrected ratio method agree closely to that of more rigorous Archie method.

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